
Ice Age 4 Continental Drift Tamil Dubbed DvD Rip 700Mb

This conclusion is consistent with the island arcs of the west coast of New Zealand, showing a thick crust along the islands and tectonized rock interpreted to be continental crust (Fig. 2). These arcs are predominantly young: the oldest age for the most active island arc is ~0.85 Ma (Higgins et al., 2015) and in the region around 1000 Ma, its age is estimated at the oldest age of the Pliocene (Lopresti et al., 2001). After recognition of a large continental crust in southwest Pacific Ocean, it is important to understand the factors that gave rise to the extension of this crust. A close examination of the geology of southwest Pacific Ocean shows that there are a series of major tectonic episodes: from ~300 Ma on the mid-ocean ridge to ~830 Ma, when the supercontinent Gondwana (with protocontinent Rodinia in the middle) started to break apart into Rodinia, Pannotia and Australia (e.g., Matthews and Richards, 1998). After that, Zealandia was situated in the southwest Pacific as a result of compressive extension during supercontinent assembly (e.g., DeMicco et al., 2003; Watson et al., 2004; Hopkinson et al., 2008). There are so many tectonic events that it is difficult to discover the critical ones when interpreting the data. As shown in Figure 4, most of the crustal thickness has accumulated through extension, being thinned down as the crust is being lengthened and thinned (e.g., Brand et al., 2001). Therefore, at least some of the extension that formed Zealandia has been from crustal extension, whereas at least part of the sea floor that was formerly above continental crust has been formed through oceanic crustal extension or crustal shortening. And, the convergence from Rodinia to Zealandia that we are now observing is a result of extension, not tectonic compression due to subduction of a large continental margin. A wider recognition of the continental margin as the basis for extension, rather than as an active plate boundary, will assist our understanding of this important process.

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The unbroken crust of the ocean basins also shows that continental crust can be broken apart as a result of long-term tectonic processes, for example, involving collision or subduction, tectonic extension, or gravitational collapse. The three main types of continental crust are continent-ocean crust (the crust of

continents that have remained separate from the ocean), continent-continent crust (the crust of continental fragments), and island-continent crust (the crust of oceanic islands). The mechanisms of formation of these three classes of crust have been discussed at length elsewhere (e.g., Scaffet et al., 1996). Island-continent crust is the result of gravitational

collapse, and island-continent collisions (Fig. 7). Continental-continent crust is a result of subduction-collision crust, and the final creation of continent-continent crust is the result of the combined effects of some or all of the above processes (Fig.

Even with an allowance for ongoing crustal thinning, the average crustal thickness of Zealandia is approximately 3.4 Mkm, and the crustal

thickness is as much as 5.0 Mkm at some places (Mortimer and Campbell, 2014). The

pattern of deformation that created Zealandia is not expected to have a large role in constraining current plate-tectonic modelling or future paleogeography.

However, a thinned continental crust does affect the total area of continental crust and hence may play a role in the effects of future

climate change. In this paper we systematically review the geology, geophysics, and geodesy of the southwestern Pacific, from the Auckland Islands in the north, to the New Caledonia Trough in the south. We review the geology and geophysics of southwest Pacific crustal margins such as the New Caledonia Trough, the Lau Fracture Zone, and the New Zealand

Trough, and continue on to the west to the Australian Murray-Darling Basin and the East Scotia Arc. This work is a culmination of previous research and new data sets and the synthesis of this data allows us to provide the first detailed geologically stable picture of the spatial extent of the southwest Pacific continental crust (Fig. 1). This supports the hypothesis that the western part of the

southwest Pacific
continental margin is
separate from the
eastern part; this is
the basis for the
scientific case for
Zealandia.
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